[Journal of Atmospheric and Solar-Terrestrial Physics](http://dx.doi.org/10.1016/j.jastp.2015.12.006) ∎ (∎∎∎∎) ∎∎∎–∎∎∎

Contents lists available at [ScienceDirect](www.sciencedirect.com/science/journal/13646826)

Journal of Atmospheric and Solar-Terrestrial Physics

journal homepage: <www.elsevier.com/locate/jastp>

Catalogue of electron precipitation events as observed in the long-duration cosmic ray balloon experiment

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article info

Article history: Received 21 October 2015 Received in revised form 8 December 2015 Accepted 10 December 2015

Keywords: Atmosphere Cosmic rays Balloon experiment Electron precipitation Geomagnetic disturbances

1. Introduction

From the middle of 1957 till present the P.N. Lebedev Physical Institute of the Russian Academy of Sciences (Laboratory of Solar and Cosmic ray physics-Dolgoprudny Scientific Station) has carried out the regular balloon measurements of charged particle fluxes in the atmosphere from the ground level up to 30–35 km above the sea level. The measurements are performed at polar (northern and southern) and middle latitudes (including six latitude surveys in1962–1987; [Bazilevskaya and Svirzhevskaya \(1998\)\)](#page-17-0). More than 85,000 measurements of cosmic ray fluxes in the atmosphere have been performed using the Geiger counters. The main goals of observations are the investigations of galactic cosmic ray modulation in the heliosphere, solar cosmic ray generation and propagation, precipitation of energetic electrons from the Earth's magnetosphere and the role of charged particles in the atmospheric processes ([Charakhch](#page-17-0)'yan, 1964; [Bazilevskaya et al., 1991;](#page-17-0) [Bazi](#page-17-0)[levskaya and Svirzhevskaya, 1998](#page-17-0); [Stozhkov et al., 2001](#page-18-0), [2009](#page-18-0)).

In this paper we focus on the magnetospheric electron precipitation into the atmosphere which is important process of the outer radiation belt depletion (e.g., [Horne and Thorne, 2003\)](#page-17-0). In the stratospheric experiment we deal with $E > 200$ keV electron flux. The magnetosphere reacts to the disturbed solar wind with changes of its configuration and violation of the magnetopause. The outer radiation belt populated by energetic electrons gets additional electrons both from the solar wind and the ionosphere.

<http://dx.doi.org/10.1016/j.jastp.2015.12.006> 1364-6826/& 2015 Elsevier Ltd. All rights reserved.

ABSTRACT

Since the International Geophysical Year (1957), the Lebedev Physical Institute performs the regular measurements of charged particle fluxes in the Earth's atmosphere (from the ground level up to 30– 35 km) at several latitudes. The unique experimental data base obtained during 58 years of cosmic rays observations in the atmosphere allows to investigate temporal, spatial and energetic characteristics of galactic and solar cosmic rays as well as the role of charged particles in the atmospheric processes. Analysis of this data base also revealed a special class of numerous events caused by energetic electron precipitation recorded in the atmosphere at polar latitudes. In this paper we present Catalogue of electron precipitation events observed in the polar atmosphere during 1961–2014 and briefly outline the previous results of this data set analysis.

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The electrons are then subject of two competing processes-betatron or stochastic acceleration and losses via escaping back to space or into the atmosphere ([Reeves et al., 2003](#page-18-0)). Currently accepted rapid loss mechanisms include magnetopause shadowing and/or outward diffusion, and precipitation to the atmosphere due to wave-particle interactions ([Ukhorskiy et al., 2015\)](#page-18-0).

Satellites situated in the magnetosphere observe an enhanced variability of high-energy electron fluxes not only because of real particle acceleration and loss but also because of particle spatial and energy redistribution. Therefore from the satellite observations it is not always easy to assess real electron precipitation. In the atmosphere we avoid this ambiguity. It should be noted that solar wind-magnetosphere coupling leading to electron precipitation remains still an open area for research ([Blum et al., 2015;](#page-17-0) [Clilverd et al., 2010;](#page-17-0) [Sandanger et al., 2009\)](#page-18-0).

Precipitation is a consequence of the electron scattering into the loss cone which results from the nonlinear wave-particle interaction. ULF and VLF waves are generated in the magnetosphere by interplanetary disturbances [\(Reeves et al., 2003\)](#page-18-0). Whistler mode chorus, hiss and electromagnetic ion cyclotron waves (EMIC) can both accelerate electrons to higher energies and scatter electrons into the loss cone through resonant pitch angle interaction. In the course of the bounce and drift motions an electron passes to various regions and can be accelerated or lost. Numerous works are devoted to modeling and observation of such effects, to mention just a few: [Carson et al., 2013](#page-17-0); [Clilverd et al., 2015;](#page-17-0) [Ker](#page-17-0)[sten et al., 2011;](#page-17-0) [Kubota et al., 2015;](#page-17-0) [Li et al., 2013;](#page-17-0) [Lorentzen et al.,](#page-17-0) [2000](#page-17-0); [Meredith et al., 2002;](#page-18-0) O'[Brien et al., 2003,](#page-18-0) [Wang et al., 2014.](#page-18-0)

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Even a brief outline shows extremely complicated dynamics of magnetospheric electrons.

Study of electron precipitations is important both for fundamental science and human practice activity. Large electron fluxes affect orbital satellites causing an electrostatic discharge (ESD) anomaly. Cumulative effect of radiation damages is one of the most important factors limiting the lifetime of a spacecraft. A single energetic electron can introduce errors into memory chips and other electronic devices, known as a single-event upset (SEU). SEUs can lead to corruption of data in memory chips and to phantom commands [\(Horne, 2002\)](#page-17-0). Additional ionization created by the electron fluxes violates the radio wave propagation in the ionosphere ([Clilverd et al., 2010](#page-17-0)). Electron precipitation contributes to the production of odd nitrogen NO_x and odd hydrogen HO_x through ion-molecular reactions in the upper atmosphere. Both NO_x and HO_x can destroy odd oxygen through catalytic reactions, and hence play an important role in the ozone balance of the middle atmosphere ([Clilverd et al., 2009;](#page-17-0) [Turunen et al., 2009;](#page-18-0) [Krivolutsky and Repnev, 2012\)](#page-17-0). During the last decades enormous efforts have been undertaken for understanding and forecasting of electron precipitation. Dynamics of energetic electrons in the magnetosphere has been observed at many spacecraft missions including GOES, POES, CRRES, SAMPEX, Polar, Cluster and others. A dedicated pair of satellites – Van Allen Probes – was launched in 2012 [\(Spence et al., 2013\)](#page-18-0). Special balloon campaigns MAXIS, MINIS, BARREL have been organized with the aim of simultaneous observation of electron fluxes in space and in the atmosphere ([Millan et al., 2002](#page-18-0), [2007;](#page-18-0) [Comess et al., 2013](#page-17-0); [Sample, 2013;](#page-18-0) [Woodger et al., 2015](#page-18-0)). Observations proved that the electron flux can vary by several orders of magnitude on time scales as short as a few minutes (e.g., [Blake et al., 1996;](#page-17-0) [Nakamura et al., 2000;](#page-18-0) [Blum](#page-17-0) [et al., 2015](#page-17-0)).

This paper presents the Catalogue of events of high-energy magnetospheric electron precipitation recorded by the cosmic ray group from the Lebedev Physical Institute during more than half century of cosmic ray observations in the stratosphere. Results of simulations of the energetic electron flux propagation through the Earth's atmosphere are described. These results were used for evaluation of characteristics of precipitating electron fluxes given in the Catalogue.

2. Balloon cosmic ray experiment in the Earth's atmosphere

Since the International Geophysical Year (1957), the Lebedev Physical Institute performs the regular measurements of cosmic rays in the atmosphere ([Charakhch](#page-17-0)'yan, 1964; [Bazilevskaya et al.,](#page-17-0) [1991;](#page-17-0) [Bazilevskaya and Svirzhevskaya, 1998](#page-17-0); [Stozhkov et al., 2001,](#page-18-0) [2009\)](#page-18-0). Observations are taken with light balloons at several latitudes almost every day. In the frame of this program, the electron precipitation events (EPEs) are detected at polar latitudes. Information on polar stations of cosmic ray measurements in the stratosphere where these events were recorded is given in Table 1.

Balloon launching time during many year observations remained in the limits of 8–11 LT and \sim 13–18 LT. A typical flight lasts about 1.5 h and a balloon usually reaches altitudes where precipitation can be observed. The balloons do not remain long at high altitudes where EPEs may be observed, i.e., we usually do not record an EPE start and end as it would be possible during longlasting balloon flights (e.g., [Lazutin et al., 1982;](#page-17-0) [Parks et al., 1993;](#page-18-0) [Millan et al., 2002\)](#page-18-0). In spite of the fact that observational time at Murmansk region and at Mirny observatory was close, almost all of EPEs were recorded at Murmansk region. This is because Mirny is located in the polar cap, mainly at the open geomagnetic field lines [\(Makhmutov et al., 2002\)](#page-17-0).

Simultaneous observations of the EPEs at the different locations are very important from the point of view of estimation of longitudinal extension of energetic electron precipitation region. Such simultaneous balloons being at the same atmospheric altitude at any two stations during the EPE are rather rare occasion. Table 2 presents a comparative statistics of few electron precipitation events observed simultaneously at pair of stations Olenya-Norilsk and Olenya – Tixie Bay.

In spite of small statistics of events recorded at Norilsk and Tixie Bay, it is possible to conclude that (1) the events recorded at Norilsk in \sim 30% of cases were also recorded at Olenya, and contrary, the events observed at Olenya in \sim 23% cases were also recorded at Norilsk; (2) in \sim 50% of cases the events recorded at Tixie Bay also were observed at Olenya, but in \sim 20% of cases the electron precipitation events at Olenya were seen at Tixie Bay. The fact of simultaneous observations of EPEs at well separated locations means that the energetic electron precipitation region

Table 1

List of high-latitude stations of stratospheric cosmic ray monitoring $(R_c$ geomagnetic cutoff rigidity).

Table 2

Comparative statistics of EPEs recorded at Olenya, Norilsk and Tixie Bay.

sometimes is widely extending over a longitude $\sim 80^\circ$. Also we point out that even though the observation of same event at pair stations is not absolutely at one time, the significant fluxes of very energetic photons were recorded at both sites.

3. Selection of the electron precipitation events

The probe for measuring of cosmic rays consists of two Geiger tubes arranged as a telescope with a 7 mm (2 g cm $^{-2}$) thick Al filter inserted between the counters. A single counter is sensitive to electrons $(E > 200 \text{ keV})$, protons $(E > 5 \text{ MeV})$, and X-rays $(E>20 \text{ keV})$, efficiency $\approx 1\%$), and a telescope records electrons $(E>5$ MeV) and protons $(E>30$ MeV). Efficiency of the charged particles recording is close to 100%. A radio pulse caused by a charged particle passing through a counter or a telescope and information on the residual air pressure is transmitted to the ground-level receiver (more details are in [Charakhch](#page-17-0)'yan, 1964; [Bazilevskaya et al., 1991](#page-17-0); [Bazilevskaya and Svirzhevskaya, 1998;](#page-17-0) [Stozhkov et al., 2001](#page-18-0), [2009\)](#page-18-0).

We identified an EPE if we observe an enhancement of the count rate of a single counter by $>$ 30%, but not of a telescope. The enhancement observed at altitude more than 20 km above the sea level lasts at least 10 min. Actually, precipitating electrons are absorbed at levels of a few $g \text{ cm}^{-2}$ in the upper atmosphere. However, they generate X-rays that can penetrate rather deep into the atmosphere and may be detected by a single counter, which is sensitive to X-rays, but not by a telescope. In \sim 75% of EPEs recorded, X-rays propagate in the atmosphere down to \sim 25 km, one third of them being registered at altitudes above 30 km. Thus, we mainly deal with precipitation of electrons with energies above several hundreds of keV ([Makhmutov et al., 2001](#page-17-0)).

Figs. 1 and [2](#page-3-0) give some examples of EPEs recorded in the atmosphere. Note that a telescope count rate does not increase while that of a single counter goes up. It is also seen that significant time variations in the X-ray flux may sometimes be observed against the quiet flux of charged particles detected by a telescope. In the absence of such variations, we assume that the electron flux at the atmospheric boundary is stable and the change in the X-ray flux is caused by the X-ray absorption in the atmosphere. Although the observations hint at variety of EPE types observed we do not make any selection by event type of the EPEs for this publication.

Vertical arrows ([Fig. 2](#page-3-0), right) show atmospheric depth levels (X_{max}) where count rates of the CR probe begin increase during the events presented at left panel of [Fig. 2](#page-3-0). The values of X_{max} were determined for whole set of precipitation events. Analysis of the X_{max} distribution allows to conclude that majority of EPEs were observed at the atmospheric depths $X < 30$ g cm⁻² (altitudes $H > 24$ km), but in several events X-rays propagate down to $X>50$ g cm⁻² (altitudes $H<20$ km). Note that the bremsstrahlung photons with initial energy E_{ph} = 0.3, 1.5 and 3 MeV have attenuation lengths in air $X=9.4$, 19.4 and 28.9 g cm⁻², correspondingly. Such energetic photons could be produced by hundreds keV – few MeV precipitating electrons at the top of the Earth's atmosphere.

To evaluate the photon absorption spectrum (ΔN values of the omnidirectional counter at different levels of X) for the events the background cosmic ray curve (curve 3 in [Fig. 2\)](#page-3-0) was subtracted from the curves 1 and 2 obtained during these EPEs. Results of such subtraction were shown in [Fig. 2](#page-3-0) (right panel). Line 1 and 2 show the best fit of the data in the form $\Delta N = N \cdot \exp(-X/X_0)$. Such approximation of the data allows to evaluate photon absorption spectra recorded in the atmosphere in the form $J_{ph(>20 \text{ keV})}(X) = A_{ph} \cdot \exp(-X/X_0)$. We determined such photon absorption spectrum for each electron precipitation event observed in the atmosphere. These parameters of the recorded photon absorption spectra, A_{ph} and X_0 , were used for evaluation of the energy spectrum of precipitating electrons at the top of the atmosphere for each event as it is described below.

4. Estimation of the precipitating electron flux and energy spectra

As mentioned before, during the EPEs a flux of secondary bremsstrahlung photons is generated by a precipitating electron flux at the top of the atmosphere. These photons can propagate down to altitudes of 15–20 km in the atmosphere and can be registered by a balloon-borne probe. The main characteristic of an electron precipitation event observed in the atmosphere is a photon absorption spectrum ([Fig. 2](#page-3-0), right panel; more details are in [Makhmutov et al. \(2003a](#page-17-0), [2003b,](#page-17-0) [2003c\)](#page-17-0)). In case of smooth increasing of the X-ray fluxes with altitude we can assume that the flux of precipitating electrons at the atmospheric boundary during the measurements is constant. In order to estimate the energy spectrum of the incident electrons we have simulated the electron transport in the atmosphere which includes the production of the secondary photons and their transport through the atmosphere,

Fig. 1. Three minute averages of count rates of single counters during (1) EPEs and (2) quiet conditions, and (3) telescopes versus residual atmospheric depth (X). Count rates of telescopes are multiplied by 5. Left panel – the EPE of September 28, 1997. Right panel – EPE of October 9, 1998.

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Fig. 2. Left: three minute averages of count rates of single counters during two EPEs (1: 29.09.1999, 08:19-09:10 UT; 2: 05.05.2000, 08:32-09:29 UT) and 3: prior to the EPEs the cosmic ray background during quiet conditions versus residual atmospheric depth. Right: > 20 keV photon absorption spectra (ΔN versus atmospheric depth (X), where ΔN are three minutes averages of the excess of single counter rates above the background) during the EPEs 1 and 2 shown at the right panel. Line 1 and 2 show the best fit of the data in the form $\Delta N = N \cdot \exp(-X/X_0)$.

similarly to other works (e.g. [Berger and Seltzer, 1972](#page-17-0); [Berger et al.,](#page-17-0) [1974](#page-17-0); [Lazutin, 1979](#page-17-0); [Kalinina et al., 1988](#page-17-0)). However, our consideration was specially directed to simulation as close as possible to conditions of our balloon measurements: (a) we chose $E = 20$ keV as energy threshold of photon registration by the STS-6 Geiger counter used in our CR probe; (b) the results of calculations were summarized for the range of atmospheric depth (0–60 g cm $^{-2}$), where the EPEs mostly were recorded in the atmosphere. In calculations we have used the Monte-Carlo ATMO-COSMICS (PLANETOCOSMICS) code based on Geant4 ([Agostinelli](#page-17-0) [et al., 2003;](#page-17-0) [Desorgher et al., 2003](#page-17-0); [Desorgher, 2004\)](#page-17-0). This code simulates the hadronic and electromagnetic interactions of energetic particles at energies $E < 100$ GeV in the Earth's atmosphere (atmospheric model MSISE90 or NRLMSISE00 or TABLE; [Desorgher](#page-17-0) [2004](#page-17-0); [Picone et al., 2002\)](#page-18-0). It computes the resulting flux of atmospheric shower particles at different altitudes (or atmospheric depths), the energy deposited in the atmosphere versus altitude, and the production of cosmogenic nuclides. The electromagnetic shower is simulated by taking into account the following processes: bremsstrahlung, energy loss by ionization, multiple scattering, pair production, Compton scattering, and photoelectric effect. First results of simulations of monoenergetic electron fluxes propagation in the atmosphere were presented in [Makhmutov](#page-17-0) [et al. \(2003a\)](#page-17-0). They confirm good agreement with those of [Berger](#page-17-0) [and Seltzer \(1972\)](#page-17-0).

To estimate the energy spectrum of precipitating electrons during EPEs, we made an assumption that the primary flux of precipitating electrons at the top of the atmosphere can be expressed by an exponential spectrum, $J_e(E) = A_e \cdot \exp(-E/E_0)$ (e.g. [Comess et al., 2013](#page-17-0); [Lazutin, 1979,](#page-17-0) [1986](#page-17-0); [Millan et al., 2007\)](#page-18-0), with characteristic energy E_0 in the range 10 keV to 1 MeV. To study the transport of primary electron flux in the atmosphere we have calculated the electron and photon energy spectra at different atmospheric depth levels. As an example in Fig. 3 we show the evolution of the energy spectrum $J_e(E) = 1 \cdot \exp(-E/300 \text{ keV})$ of incident electrons on the top of the atmosphere at selected atmospheric depth levels $X=0.05$, 0.5 and 1 g cm⁻² (64.5, 49.9 and 45.5 km, correspondingly).

Note very significant reduction (by factor $\sim 10^{4-6}$) of primary electron fluxes already at altitude \sim 50 km (X=0.49 g cm⁻²).

Fig. 4 demonstrates the secondary photon energy spectra at several atmospheric levels. As can be seen the photon flux (e.g., at energy $E = 100$ keV) at altitude 24 km decreased only by factor 15 regarding to the photon flux level at 70 km. In comparison, the

Fig. 3. Electron energy spectra at several atmospheric levels $X=0.05$ (curve 1), 0.5 (curve 2) and 1 g cm⁻² (curve 3). Electron flux $J(E)$ in arbitrary units (a.u.).

Fig. 4. Photon energy spectra at several atmospheric levels $X=0.05$ g cm⁻² (altitude \sim 70 km; curve 1), $X = 15$ g cm⁻² (altitude = 28.3 km; curve 2) and $X=30$ g cm⁻² (altitude = 23.9 km; curve 3). Photon flux $J_{ph}(E)$ in arbitrary units (a. u.).

secondary electron flux reduced by factor \sim 10⁷ at those altitudes.

Finally we have computed the number of secondary photons with $E > 20$ keV at different atmospheric depths (in the range $X=0.05-50$ g cm⁻²) resulting from precipitation of primary electron flux with different exponential parameter E_0 . For each E_0 we have obtained a photon absorption spectrum, i.e. photon number

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Fig. 5. Simulation results: left panel – the dependence of the parameter X_0 of the photon absorption spectrum in the atmosphere on the characteristic energy E_0 of the primary electron spectrum in the form of $J_e(E) = A_e \cdot \exp(-E/E_0)$; right panel – the correspondence between the coefficient A_{ph} of the photon spectrum and characteristic energy E_0 of the electron spectrum.

versus atmospheric depth, that we fit with an exponential law as $J_{ph(>20 \text{ keV})}(X) = A_{ph} \cdot \exp(-X/X_0)$. Fig. 5 (left) shows the parameter X_0 of the photon absorption spectrum as a function of the characteristic energy E_0 of the primary electron spectrum. Using the results shown in Fig. 5, we can deduce the spectral parameter E_0 for a specific EPE from the parameters X_0 and J_0 of the observed absorption spectrum ([Makhmutov et al., 2003a](#page-17-0), [2006\)](#page-18-0). For retrieving these parameters we have used the 3-min averaged data and fitted the experimental X-ray absorption spectrum by an exponent. Thereby we did not account for fast temporal variations of the photon flux. In the case of several separate count rate enhancements during one balloon flight we have treated only that at the highest altitudes (therefore some powerful event may be missed).

This procedure was applied to each EPE and the results are presented in the [Table A1](#page-5-0) as the Catalogue of EPEs recorded by our group in the course of the long-term charged particle fluxes monitoring in the stratosphere since 1957 (see Attachment). Some precipitations take place during intrusion into magnetosphere of solar energetic particles. Only one event was discussed in ([Bazi](#page-17-0)[levskaya et al., 2002\)](#page-17-0). Such events are not included in this data set and will be analyzed later.

For each precipitation event in the Catalogue the following information is given:

- event number,
- date (dd mm yyyy),
- time interval of the EPE observation (Start Time, T_0 End Time, T_e ; hh mm hh mm, UT),
- $-A_e$ parameter of the flux of incident electrons in cm^{-2} s⁻¹ keV⁻¹,
- E_0 characteristic energy of electron spectra in keV,
- $J(E > 20 \text{ keV})$ integral flux of incident $> 20 \text{ keV}$ electrons in $\rm cm^{-2} \, s^{-1}.$

It should be noted that parameters of the energy spectra (A_e) and E_0) as well as the value of >20 keV electron flux (*J* $(E>20 \text{ keV})$) were obtained under assumption that the incident electron flux did not change during the time interval T_0-T_e . In reality the electron flux can fluctuate rapidly on the time scale of several minutes [\(Blake et al., 1996](#page-17-0); [Nakamura et al., 2000;](#page-18-0) [Blum](#page-17-0) [et al., 2015\)](#page-17-0). Therefore, the parameters A_e , E_0 and $J(E>20 \text{ keV})$ should be considered as a first approximation or an estimation.

5. Conclusion

The time series of EPEs in the stratosphere presented in the Catalogue, were investigated in a number of works (e.g., [Bazilevskaya and Svirzhevskaya, 1998;](#page-17-0) [Makhmutov et al., 2001,](#page-17-0) [2002,](#page-17-0) [2005,](#page-17-0) [2012](#page-18-0), and references therein). The obtained results of the analysis could be briefly outlined as follow.

- 1. On the long-term scale, the EPE yearly rate reaches maximum at the descending phase of solar activity cycle and correlates with occurrence of corotating high-speed solar wind streams. Apparently, the CME-related transient solar wind disturbances typical for solar maximum conditions are not very effective for EPE production in the Earth's atmosphere ([Makhmutov, et al.,](#page-17-0) [2003a\)](#page-17-0).
- 2. The monthly numbers of the EPEs recorded at Olenya station (Murmansk region) show the semiannual variation with two maxima. The first one is in April and the second one is rather extended covering August-October period. We believe that the first peak is in accordance with the expectation of Russel-McPherron effect. A second peak is complex and probably due to the superposition of axial, equinoctial and Russel-McPherron effects [\(Makhmutov et al., 2003c,](#page-17-0) [2005\)](#page-17-0).
- 3. On the day-to-day scale, EPEs most probably occur under southward B_z , two days after SSC, one day after maximum of the IMF B_z strength. On the EPE day, maximum values of solar wind velocity, K_p , AE, and D_{st} indices are observed. EPEs most probably happen during significant increase in the relativistic electron flux at geostationary orbit ([Makhmutov et al., 2003c](#page-17-0)).

Our previous findings did not consider properties of precipitating electrons. The present Catalogue will be used for study of the relations between the parameters of the electron spectra and concomitant phenomena with the aim of better understanding of the underlying physics. An area of electron precipitation research is still very wide. More information is needed about dynamics of the energy spectra, temporal variability of precipitation, relative contribution of various mechanisms in the radiation belt depletion, depending on the type of geomagnetic disturbance. Hopefully, the present Catalogue will be useful also for modeling of atmospheric effects such as additional ionization production of NO_x and HO_x during almost five cycles of solar activity.

Acknowledgement

The cosmic ray monitoring in the atmosphere is partly supported by Russian Foundation for Basic Research Grants no.14-02- 00905a, 13-02-00585, 13-02-00931, 15-02-10070k, 16-02-00100 and by the Program "High Energy Physics and Neutrino Astrophysics" of the Russian Academy of Sciences. The Catalogue of EPE was compiled during implementation of the ISSI team project

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Table A.1

Catalogue of Electron Precipitation Events recorded in the long-term cosmic ray experiment in the stratosphere at the polar stations Murmansk region (Olenya, Apatity), Norilsk, Mirny and Tixie Bay. The 3-min averages of data were used. For each event the following information is given: event number date (day, month, year) time interval of
the EPE observation (T_o-T_e, Start and End Tim of electron spectra in keV J(E > 20 keV) – integral flux of incident $\,$ > 20 keV electrons in cm⁻² s⁻¹.

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Table A.1 (continued)

Table A.1 (continued)

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Table A.1 (continued)

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Table A.1 (continued)

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Table A.1 (continued)

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Table A.1 (continued)

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Table A.1 (continued)

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Table A.1 (continued)

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Table A.1 (continued)

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Table A.1 (continued)

Norilsk

Mirny

Tixie Bay

Table A.1 (continued)

2013–2015 and by the Program "Specification of ionization sources affecting atmospheric processes" leaded by Irina Mironova.

Appendix

See [Table A1](#page-5-0)

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